



1 **Reconstructing Sea Level Variability at the Jeodo Ocean Research**
2 **Station (1993–2023) Using Artificial Intelligence, Machine Learning,**
3 **and Reanalysis Integration**

4

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11



Abstract

This study presents a comprehensive approach for reconstructing a high-quality, continuous monthly sea level time series at the Jeodo Ocean Research Station (IORS) from 1993 to 2023 using advanced artificial intelligence (AI) and machine learning (ML) models. After applying quality control to the in-situ KIOST data, including inverse barometric effect correction, 3σ filtering, and a 75% data coverage threshold, we validated trends using nearby PSMSL tide gauges and four ocean reanalysis datasets (CMEMS, GLORYS, ORAS5, HYCOM). The trend analysis showed a higher rate of sea level rise from in-situ data (4.94 mm/yr, Oct 2003–Dec 2023) compared to satellite and model-based estimates (e.g., CMEMS: 3.53 mm/yr, Jan 1993–Dec 2023), suggesting localized sea level rise in the East China Sea. Initial gap-filling used statistical models such as harmonic regression and regression-based climatology. A blended approach combining climatology and trend components achieved the best accuracy (RMSE ~ 0.056 m, $R^2 = 0.688$). We then implemented various AI/ML models through an Iterative Imputer framework. Ensemble models (e.g., XGBoost) performed perfectly after 2003 but did not generalize well before 2004. Deep learning models like LSTM and GRU effectively captured seasonal and nonlinear patterns post-2003, with LSTM achieving RMSE = 0.023 m and $R^2 = 0.95$. Time series models Prophet and SARIMA-SIN successfully reconstructed the full time series, with SARIMA-SIN estimating the highest trend (5.61 mm/yr). Multiple linear regression using reanalysis data served as a baseline, but AI/ML models outperformed it in both accuracy and generalization. This study provides a reproducible, interpretable, and physically consistent framework for reconstructing sea level variability in semi-enclosed coastal seas.

Keywords: Sea Level Variability, Jeodo Ocean Research Station, AI-based Imputation,



35 Timeseries Reconstruction, East China Sea



36 **1. Introduction**

37 Long-term sea level observations are critical for monitoring regional oceanographic and
38 climatic changes, particularly in coastal and marginal seas where variability is amplified by
39 topographic and atmospheric forcing (Hamlington et al., 2020; Cazenave and Moreira, 2022).
40 The Ieodo Ocean Research Station, located at the intersection of the East China Sea (ECS) and
41 the Southern East Sea (Sea of Japan), plays a pivotal role in observing regional sea level
42 variability and marine environmental conditions (Han, 2020; Byun et al., 2021). As a
43 strategically important in-situ platform, its various interval sea level records provide valuable
44 information for understanding the dynamics of the Kuroshio Current system, East Asian
45 monsoon variability, and climate change-driven sea level rise (Ha et al., 2019; Xu et al., 2015;
46 Chang and Oey, 2011).
47 However, long-term in-situ observations are often interrupted by equipment failure,
48 maintenance issues, or extreme weather conditions such as typhoons (Adebisi et al., 2021).
49 These disruptions lead to temporal gaps in the observational records, which hinder the detection
50 of trends, reconstruction of seasonal cycles, and validation of satellite and model-based
51 products (Beguiria et al., 2019). In regions like the ECS, where strong seasonal and interannual
52 signals are present, accurate and realistic imputation of missing data is essential for scientific
53 and operational applications (Lin et al., 2020; Han, 2020).



54 Various methods have been developed to address missing data in oceanographic and climate
55 time series (Kolukula and Murty, 2025; Lee et al., 2022). Traditional approaches include linear
56 interpolation, monthly climatological averages, and harmonic regression models (Schlegel et
57 al., 2019; Risien et al., 2022; Okkaoğlu et al., 2020; Arguez and Applequist, 2013). More
58 recently, advanced statistical and machine learning techniques have been proposed for gap-
59 filling, including Gaussian process regression, Kalman filtering, and neural networks (Vance
60 et al., 2022; Wenzel and Schröter, 2010; Wang, 2023). While these methods offer improved
61 flexibility and accuracy, they often require dense observations or training data, which may not
62 be feasible in long-term sparse records (Lee et al., 2022; Sarafanov et al., 2022; Park et al.,
63 2023). Interpretable and statistically robust methods remain essential for operational and
64 historical datasets such as IORS (Han and Lim, 2024; Han, 2020).

65 This study focuses on imputing, filling, and predicting gaps in the monthly sea level data at the
66 IORS over 1993-2023. We evaluate and compare three regression-based approaches: (1)
67 harmonic regression with annual and semiannual cycles plus a linear trend (Okkaoğlu et al.,
68 2020), (2) a regression-based monthly climatology model with a trend using calendar year
69 month dummy variables (Hyndman and Athanasopoulos, 2018), and (3) a pure climatology-
70 plus-trend approach based on aggregated monthly means and linear fitting (Brunetti et al.,
71 2014). Furthermore, we propose a realistic blending method that optimally combines harmonic



72 and climatological components to minimize reconstruction errors.

73 Also, we implemented an ensemble of statistical and machine learning (ML) models to

74 reconstruct missing monthly values at the IORS (Tong and Li, 2025; Bochow et al., 2025).

75 These models were selected to span a broad range of algorithmic families, including ensemble

76 tree-based methods (e.g., XGBoost, Random Forest, LightGBM, AdaBoost) (Niazkar et al.,

77 2024; Wu et al., 2024; Gan et al., 2021; Xiao et al., 2019), regularized linear models (Lasso,

78 Ridge) (Pan et al., 2025), proximity-based models (K-Nearest Neighbors) (Latif et al., 2024),

79 and neural networks (LSTM, GRU) (Sun et al., 2020; Tumse and Alcansoy, 2025). Each model

80 was trained to impute missing values using only observed sea level data from the original

81 dataset, applying the IterativeImputer framework with consistent hyperparameters for

82 comparability (Ramirez et al., 2023). Complementing these were two timeseries specific

83 models: Facebook’s Prophet (Elneel et al., 2024), which decomposes time series into seasonal

84 and trend components, and SARIMA (Sun et al., 2020), which captures both non-seasonal and

85 seasonal autocorrelation structures. All models were evaluated using root mean square error

86 (RMSE) and coefficient of determination (R^2), calculated against observed (non-missing)

87 values (Boursalie et al., 2022; Siddig et al., 2021).

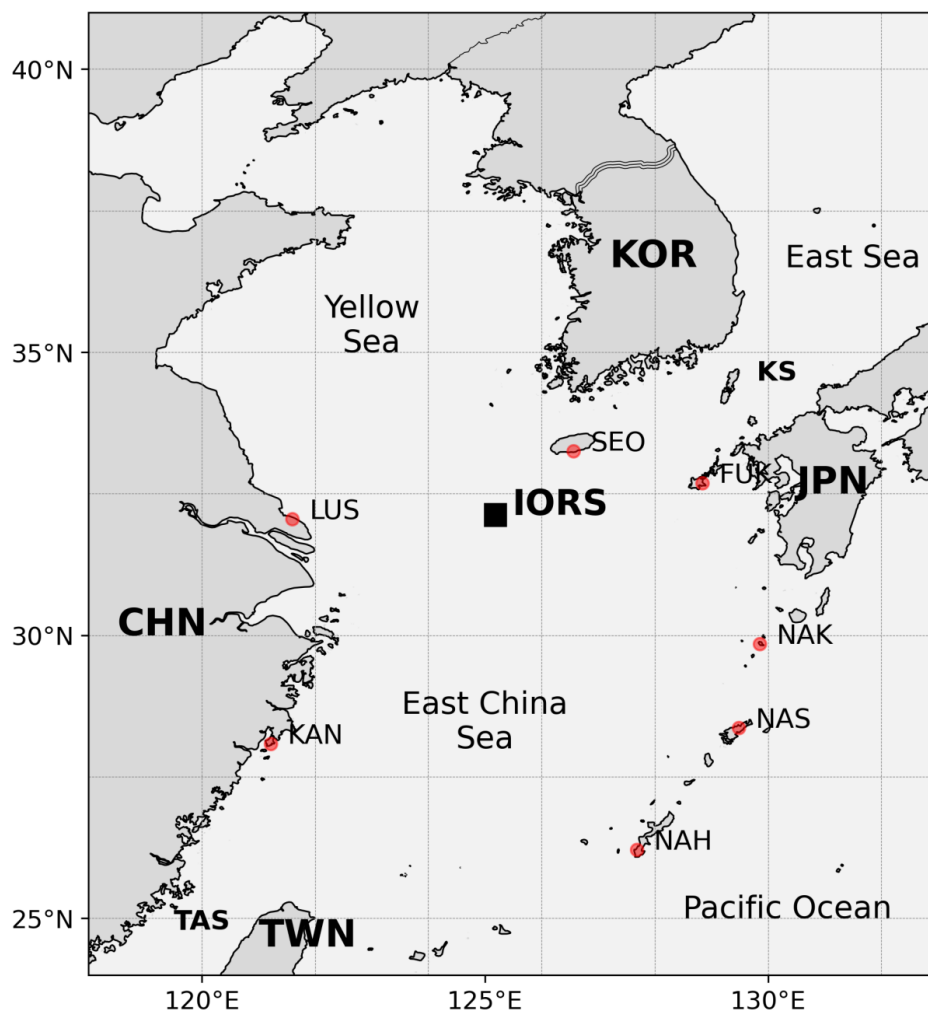
88 To contextualize the reconstructed series, we further compared the IORS observations with

89 monthly sea level data from four global ocean reanalysis products—Copernicus Marine



90 Environment Monitoring Service (CMEMS), Global Ocean Physics Reanalysis System
91 (GLORYS), Ocean Reanalysis System 5 (ORAS5), and Hybrid Coordinate Ocean Model
92 (HYCOM)—over 1993–2023 (Han et al., 2024; Long et al., 2023; Jin et al., 2023b; Cummings
93 and Smedstad, 2014). Linear trends were estimated and visualized to assess the consistency
94 and fidelity of in-situ observations relative to reanalysis-based records.

95 Finally, we extended the imputation framework to a multivariate setting by incorporating these
96 reanalysis datasets as auxiliary predictors, enabling more robust and physically informed gap-
97 filling. By leveraging both statistical and AI/ML-based techniques, this study provides a
98 transparent, reproducible framework for reconstructing realistic monthly sea level time series
99 in data-sparse coastal environments.



100

101 **Figure 1. Map showing the locations of sea level observation stations used in this study,**
 102 **including the Ieodo Ocean Research Station (IORS, black square) and seven tide gauge**
 103 **stations (red circles) from the PSMSL and regional networks. The map spans the East**
 104 **China Sea, Yellow Sea, East Sea (Sea of Japan), and the western Pacific Ocean. Major**
 105 **geographic regions and oceanic features labelled, including Korea (KOR), China (CHN),**
 106 **Japan (JPN), and Taiwan (TWN); KS and TAS are Korea and Taiwan Straits, respectively.**

107



108 **2. Data and Methods**

109 2.1. Data and Preprocessing

110 This study utilizes sea level observations from the IORS, collected between October 2003 and
 111 December 2023. The data, provided by the Korea Institute of Ocean Science and Technology
 112 (KIOST), include measurements at 1-minute, 10-minute, and hourly intervals (Kiost, 2025).
 113 Data gaps occur intermittently due to sensor malfunctions, scheduled maintenance, and severe
 114 weather events.

115 To ensure physical consistency with satellite altimetry, numerical models, and reanalysis
 116 products, all sea level values were corrected for the inverse barometric effect (IBE) in Eq. (1)
 117 (Han et al., 2024). The correction was computed using collocated atmospheric pressure
 118 measurements:

$$119 \quad \eta_{\text{corrected}} = \eta_{\text{raw}} + \frac{(P - P_0)}{\rho g} \quad (1)$$

120 Where η_{raw} is the observed and uncorrected sea level (m), P is local atmospheric pressure (Pa),
 121 P_0 is the standard atmospheric pressure (101300 Pa), ρ is the assumed seawater density (1025
 122 kg/m³), and g is gravitational acceleration (9.81 m/s²).

123 A multi-tiered quality control protocol was implemented to construct reliable timeseries for
 124 subsequent analyses in Figures 2, 3 and 4:

- 125 - Quality Flag Filtering: Only values flagged as “good” based on instrument diagnostics



126 and metadata were retained (Gouldman et al., 2017) (Figure 2).

127 - 3σ statistical Filtering: Outliers exceeding ± 3 standard deviations from the local mean

128 were excluded (Oelsmann et al., 2021) (Figure 2).

129 - 75% Threshold Rule: Daily and monthly averages were retained only if at least 75%

130 of the corresponding observations passed the above filters (Epa, 2000) (Figures 3 and

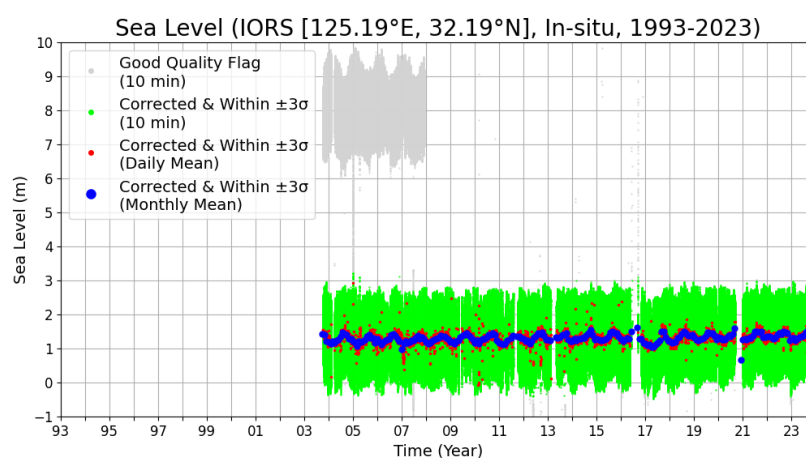
131 4).

132 The resulting dataset comprises clean daily and monthly timeseries, exhibiting a clear seasonal

133 cycle, interannual variability, and several data gaps that require reconstruction. These

134 timeseries served as the foundation for further statistical modeling and machine learning-based

135 gap-filling.



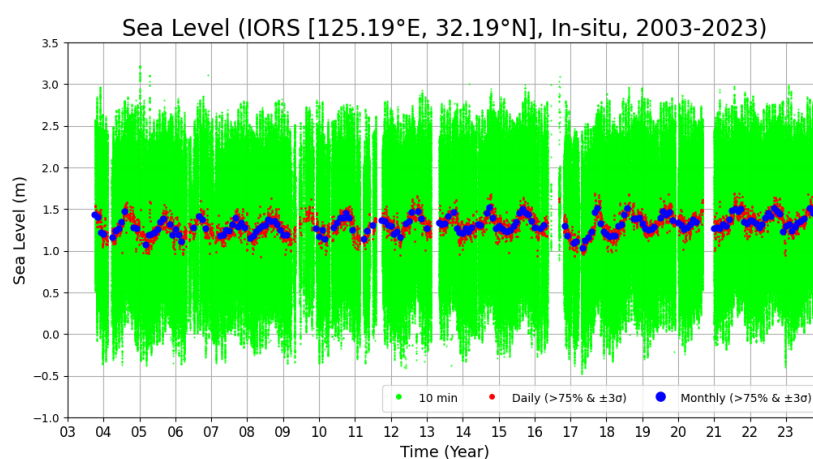
136

137 **Figure 2.** Time series of inverse barometric effect (IBE)-corrected in-situ sea level

138 **observations at the IORS from 1993 to 2023. Gray dots represent 10-minute interval data**

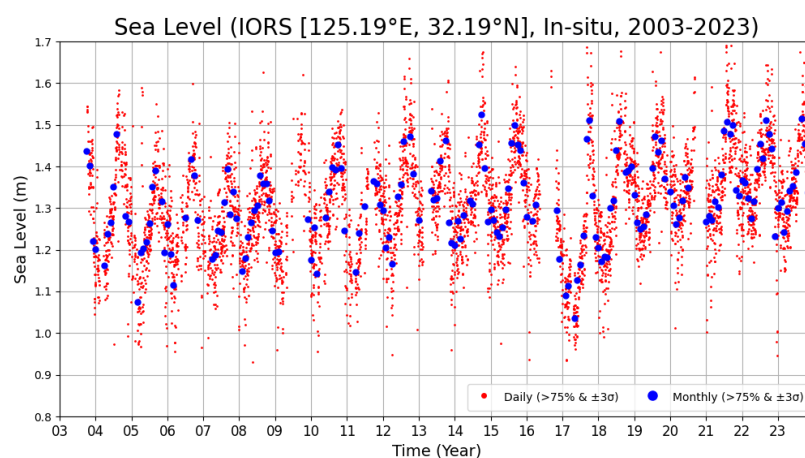


139 **flagged as good quality. Green dots indicate corrected 10-minute data within $\pm 3\sigma$. Red**
 140 **dots denote daily means within $\pm 3\sigma$, and blue dots show monthly means within $\pm 3\sigma$. Data**
 141 **gaps reflect periods of sensor malfunction or data loss.**



142

143 **Figure 3. Sea level time series at the IORS from 2003 to 2023, based on in-situ**
 144 **observations corrected for the IBE. Green dots represent 10-minute sea level data. Red**
 145 **dots show daily mean values calculated only when at least 75% of 10-minute observations**
 146 **per day are available and within ± 3 standard deviations. Blue dots denote monthly mean**
 147 **sea level, computed from daily means that meet the same 75% availability and ± 3 criteria.**



148

149 **Figure 4. Filtered daily and monthly sea level time series at the IORS from 2003 to 2023.**

150 **Red dots represent daily mean values calculated when at least 75% of 10-minute in-situ**

151 **observations are available and within ± 3 standard deviations. Blue dots indicate**

152 **monthly mean values derived from daily means that also meet the 75% data availability**

153 **and $\pm 3\sigma$ filtering criteria.**

154

155 2.2. Sea Level Reconstruction Using Regression Techniques

156 We applied three regression-based approaches to reconstruct missing values and estimate long-

157 term sea level trends:

158

159 2.2.1. Harmonic Regression with Trend (HR)

160 This method models the sea level time series as a combination of a linear trend and periodic



161 seasonal components. Specifically, it includes annual and semiannual sine and cosine terms to
 162 capture seasonal variability in Eq. (2) (Young et al., 1999):

$$163 \quad y_t = \beta_0 + \beta_1 t + \beta_2 \sin(\omega t) + \beta_3 \cos(\omega t) + \beta_4 \sin(2\omega t) + \beta_5 \cos(2\omega t) + \epsilon_t \quad (2)$$

164 where $\omega = 2\pi/12$, and t is time in months since January 1993. It captures both long-term and
 165 seasonal variability and is especially suitable for periodic signals.

166

167 2.2.2. Regression-Based Monthly Climatology with Trend (RC)

168 This method models monthly sea level as a function of a linear time trend and monthly
 169 categorical effects. Multiple linear regression uses dummy variables D_m (0 or 1) for calendar
 170 months (February through December), with January as the reference month. This allows the
 171 estimation of monthly climatology and trends simultaneously in Eq. (3). The model is
 172 expressed as (Hyndman and Athanasopoulos, 2018):

$$173 \quad y_t = \beta_0 + \beta_1 t + \beta_2 D_2 + \beta_3 D_3 + \cdots + \beta_{12} D_{12} + \epsilon_t$$

$$174 \quad = \beta_0 + \beta_1 t + \sum_{m=2}^{12} \beta_m D_m + \epsilon_t \quad (3)$$

175 Where y_t is the sea level at time t , β_0 is the intercept, β_1 captures the long-term linear trend,
 176 and β_m quantifies the deviation of each month m from January. This formulation allows for a
 177 flexible representation of the seasonal cycle without assuming a sinusoidal structure and can
 178 accommodate non-harmonic monthly anomalies while still estimating a long-term trend.



179

180 2.2.3. Pure Climatology with Trend (PC)

181 In this method, seasonal climatology is first computed by averaging the observed sea level for
182 each calendar month across all years (Jin et al., 2023a). A separate linear trend is then fitted to
183 the full-time series. The two components are summed to reconstruct the entire series in Eq. (4).
184 While simple and interpretable, this method assumes stationarity in the seasonal cycle and may
185 miss interannual variations:

$$186 \quad y_t = C_{m(t)} + (\alpha_0 + \alpha_1 t) \quad (4)$$

187 where $C_{m(t)}$ is the climatological mean for the month of t . This assumes a stationary seasonal
188 cycle.

189

190 2.2.4. Realistic Blending (RB)

191 To enhance the realism of the reconstructed sea level time series, we implemented a Realistic
192 Blending (RB) approach that combines the outputs of the Harmonic Regression with Trend
193 (HR) and the Regression-Based Monthly Climatology with Trend (RC) models in Eq. (5). The
194 final imputed value at each time step is computed as a weighted average of the two models
195 (Krinner et al., 2005):

$$196 \quad y_t^{RB} = \alpha y_t^{HB} + (1 - \alpha) y_t^{RC} \quad (5)$$



197 where $\alpha \in [0,1]$ is an optimal blending weight selected to minimize the root mean square error
198 (RMSE) between the blended estimate and the observed values. This method leverages the
199 physical interpretability of harmonic modeling and the statistical flexibility of regression-based
200 climatology, aiming to strike a balance between robustness and fidelity to observed variability.

201

202 2.3 Machine Learning and Statistical Models for Data Gap Imputation

203 To fill gaps in the monthly sea level record, we evaluate a range of machine learning and
204 statistical models that have been effective for time series gap-filling in environmental datasets.

205 Each model is described briefly below with supporting literature:

206

207 2.3.1. Extreme Gradient Boosting (XGBoost)

208 A powerful tree-based ensemble method that uses gradient boosting and inherently handles
209 missing values by learning default split directions (Niazkar et al., 2024). XGBoost has
210 demonstrated high accuracy in gap-filling environmental time series (e.g., used to impute
211 significant gaps in aerosol optical depth data successfully) (Fan et al., 2020)

212

213 2.3.2. Gradient Boosting Regressor

214 A generic gradient boosting algorithm for regression that builds an ensemble of decision trees



215 iteratively to minimize prediction error. GBR has achieved low interpolation errors in climate
216 data gap-filling, outperforming neural networks and multiple linear regression in one study of
217 meteorological time series (Otchere et al., 2022; Körner et al., 2018).

218

219 2.3.3. Random Forest (RF)

220 An ensemble of decision trees trained on bootstrapped samples, aggregating their results.
221 Random forests can capture nonlinear relationships and have robust performance for missing
222 data imputation (Wu et al., 2024; Tang and Ishwaran, 2017). In comparative experiments, tree-
223 based methods like RF rank among the top performers for reconstructing missing
224 environmental time series data (Mahabbati et al., 2021).

225

226 2.3.4. AdaBoost

227 An adaptive boosting algorithm that sequentially combines many weak learners (often shallow
228 trees) to improve prediction accuracy. The AdaBoost framework, while originally devised for
229 classification, has a regression variant that has been applied to time series gap-filling, for
230 example, by ensembling multiple neural-network base learners to impute missing traffic flow
231 data, yielding improved accuracy and stability (Xiao et al., 2019; Shang et al., 2024).

232



233 2.3.5. LightGBM

234 LightGBM is a gradient boosting machine algorithm optimized for computational efficiency
235 and scalability, employing histogram-based feature binning and leaf-wise tree growth. It has
236 demonstrated high predictive accuracy in forecasting tasks and is widely applied in
237 hydrological gap-filling studies(Wang et al., 2025). Its ability to efficiently process large
238 datasets and natively handle missing values makes it highly suitable for gap-filling in extended
239 climate time series, while also effectively capturing the nonlinear interactions between tidal
240 dynamics and river discharge (Gan et al., 2021).

241

242 2.3.6. CatBoost (approximated via Gradient Boosting)

243 CatBoost is a gradient-boosting algorithm specifically designed to handle categorical features
244 effectively. It introduces innovative techniques such as ordered boosting and efficient encoding
245 of categorical variables, which help in reducing overfitting and improving model performance
246 (Hancock and Khoshgoftaar, 2020). In this study, due to library constraints, we approximate
247 CatBoost's functionality using scikit-learn's Gradient Boosting Regressor. While this
248 approximation does not fully replicate CatBoost's specialized handling of categorical data, it
249 allows us to implement a gradient boosting approach within our existing framework.

250



251 2.3.7. K-Nearest Neighbors

252 A non-parametric imputation approach that fills a missing value based on the values of its
253 closest neighbors in feature space. KNN is simple yet effective for gap-filling; a study
254 comparing time series imputation methods found that KNN achieved the highest reconstruction
255 accuracy for missing data in several cases (Ahn et al., 2022). The method leverages spatial or
256 temporal similarity, which can be particularly useful when neighboring station data or nearby
257 time points correlate with the missing sea level values.

258

259 2.3.8. Lasso and Ridge Regression

260 These are regularized linear regression models that impose L1 (Lasso) and L2 (Ridge) penalties,
261 respectively, to prevent overfitting. By shrinking coefficients, they provide stable estimates that
262 can be used to predict and interpolate missing values from other correlated variables.
263 Regularized regressions have been explored in gap-filling contexts to utilize correlations in
264 environmental data while avoiding multicollinearity issues (Wijesekara and Liyanage, 2023).
265 For instance, an elastic net (a combination of Lasso and Ridge) approach outperformed basic
266 ARIMA-based methods in imputing large gaps of air quality data, highlighting the value of
267 such penalized linear models.

268



269 2.3.9. Decision Tree

270 A single decision tree can be used as a regression model to estimate missing values by learning
271 piecewise constant relationships in the data. Although decision trees alone are prone to higher
272 variance, they have been applied to gap-filling and can capture nonlinear dependencies in sea
273 level time series. In practice, tree-based algorithms (and their ensembles) have been found
274 superior to many conventional interpolation techniques for climate data gaps(Zhu et al., 2023).
275 Simpler decision trees may serve as interpretable benchmarks, while often forming the building
276 blocks of more complex ensemble methods used for imputation.

277

278 2.3.10. Prophet

279 A forecasting model based on additive decompositions of trend, seasonality, and holidays
280 (developed by Facebook). Prophet is robust to missing data and irregular timing – it does not
281 require regularly spaced observations and can model gaps without explicit interpolation (Elneel
282 et al., 2024). It has been applied in environmental time series forecasting (e.g. for air quality
283 and water levels) and can be used to predict and back-fill missing monthly sea level values by
284 leveraging seasonal patterns and trends in the data.

285

286 2.3.11. SARIMA-SIN (Seasonal ARIMA with Harmonic Regression Preprocessing)



287 To enhance the capacity of traditional SARIMA models in capturing periodic signals and long-
288 term trends in oceanographic time series, we implemented a hybrid approach known as
289 SARIMA-SIN. This method integrates harmonic regression and SARIMA modelling (Wang et
290 al., 2019). In the first stage, a deterministic seasonal component is modeled using a linear
291 combination of sine and cosine basis functions representing annual and semi-annual cycles.
292 This harmonic regression accounts for the primary seasonality, after which residuals are
293 computed by subtracting the fitted seasonal signal from the original time series. This
294 preprocessing alleviates the need for SARIMA to fit both trend and periodicity simultaneously,
295 thereby improving model stability and interpretability. Compared to conventional SARIMA
296 models, SARIMA-SIN reduces overfitting and better preserves long-term variability, making
297 it particularly suitable for sea level gap filling where both seasonal dynamics and secular trends
298 are essential.

299

300 2.3.12. Neural Networks (LSTM and GRU)

301 Recurrent neural networks, particularly Long Short-Term Memory (LSTM) and Gated
302 Recurrent Unit (GRU) architectures, are well-suited for sequential data and have been
303 successfully applied to gap-filling in complex time series. These networks maintain internal
304 memory states that allow them to learn long-term dependencies, which is advantageous for



305 inferring missing segments of a sea level record. Studies have shown that LSTM models can
 306 achieve more accurate gap imputations than traditional methods for environmental sensor data
 307 (Song et al., 2020), and specialized RNN variants (e.g., GRU-D) are designed to handle
 308 sequences with missing values by learning decay rates for missing inputs (Che et al., 2018). In
 309 practice, LSTM/GRU networks can leverage the temporal patterns in sea level data (and
 310 potentially exogenous inputs) to predict missing monthly values with high fidelity.

311

312 2.4. Evaluation Metrics and Implementation

313 To assess the performance of each imputation model, we computed two widely adopted
 314 evaluation metrics:

315 2.4.1. Root Mean Square Error (RMSE)

316 RMSE quantifies the average magnitude of reconstruction error by penalizing large deviations
 317 more strongly in Eq. (6) (Montgomery et al., 2021). It is defined as

$$318 \quad \text{RMSE} = \sqrt{\frac{1}{N} \sum_{i=1}^N (y_i - \hat{y}_i)^2} \quad (6)$$

319 where y_i are the observed values and $\hat{y}_i$ are the reconstructed values at non-missing time
 320 steps.

321 2.4.2. Coefficient of Determination (R^2)

322 The coefficient of determination, R^2 , quantifies the proportion of variance in the observed data
 323 that is explained by the model in Eq. (7) (Montgomery et al., 2021). It is calculated as:



$$R^2 = 1 - \frac{\sum_{i=1}^N (y_i - \hat{y}_i)^2}{\sum_{i=1}^N (y_i - \bar{y})^2} \quad (7)$$

Where y_i , $\hat{y}_i$, $\bar{y}$, and N denote the observed values, the model-predicted values, the mean of observed values, and the number of valid (non-missing) observations.

An R^2 value of 1 indicates a perfect fit, meaning the model explains all the variance in the data. A value of 0 implies the model performs no better than simply predicting the mean, and negative values indicate that the model performs worse than the mean-based prediction. Thus, higher R^2 values reflect stronger agreement between the reconstructed and observed time series.

332

2.5. Comparison with Ocean Reanalysis Products

To contextualize in-situ observations, monthly sea level data from four global ocean reanalysis products—CMEMS, GLORYS, ORAS5, and HYCOM—were compared with the KIOST in-situ data at IORS over the period 1993–2023. All datasets were aggregated to monthly means and aligned in time for consistency. Linear trends (expressed in mm/year) were estimated using least squares regression and applied to each time series with NaN values masked. The trend slope was computed in meters per month and converted to millimeters per year by multiplying by 12,000. Each dataset was visualized on a same plot with annotated trend values to assess agreement and long-term consistency between in-situ and reanalysis records. The comparison



342 allowed us to assess model fidelity and characterize interannual-to-decadal sea level variability
343 across observational and modeled sources.

344

345 2.6. AI/ML-Guided Multivariate Imputation Using Ocean Reanalysis Data

346 To improve the fidelity of gap-filling in the KIOST sea level time series, we implemented a
347 multivariate imputation framework that integrates auxiliary predictors from four ocean
348 reanalysis products: CMEMS, GLORYS, ORAS5, and HYCOM in Table 1. Specifically, we
349 applied IterativeImputer from the scikit-learn package in Python, using multiple linear
350 regression as the underlying estimator. In this configuration, the missing values in the KIOST
351 in-situ observations were modeled as a function of the corresponding values from the reanalysis
352 datasets, assuming linear dependence among variables.

353 Each imputation process was iteratively updated over 20 cycles to ensure numerical
354 convergence and stability of the imputed results. In parallel, univariate imputation was also
355 conducted using standalone time series models—Prophet, SARIMA-SIN, LSTM, and GRU—
356 trained solely on the historical KIOST observations. These models capture temporal dynamics
357 without external predictors.

358 Model outputs were evaluated based on the consistency of reconstructed trends, visual
359 coherence in the temporal structure, and alignment with physical expectations. The integration



360 of regression-based multivariate imputation and AI/ML models forms an ensemble-based
361 approach that leverages both data-driven learning and physically informed predictors for robust
362 reconstruction of missing values.

363

364 **Table 1. Overview of the datasets used in this study, including their temporal resolutions**
365 **and time periods.**

Data Source	Temporal Resolution	Time Period
IORS	1-min, 10-min, hourly	Oct 2003 - Dec 2023
CMEMS	Daily, Monthly	Jan 1993 - Dec 2023
GLORYS	Monthly	Jan 1993 - Dec 2023
ORAS5	Monthly	Jan 1993 - Dec 2023
HYCOM	Monthly	Jan 1994 - Dec 2023

366

367 **3. Results**

368 **3.1. Comparison with Other Data**

369 A linear regression performed on the 3σ -filtered, IBE-corrected monthly mean time series
370 revealed a significant upward trend of approximately 4.94 mm/yr from October 2003 to
371 December 2023 and 5.43 mm/yr from January 2004 to December 2023. This value is consistent
372 with other in-situ coastal tide gauge trends in the ECS and reflects regional steric and mass-



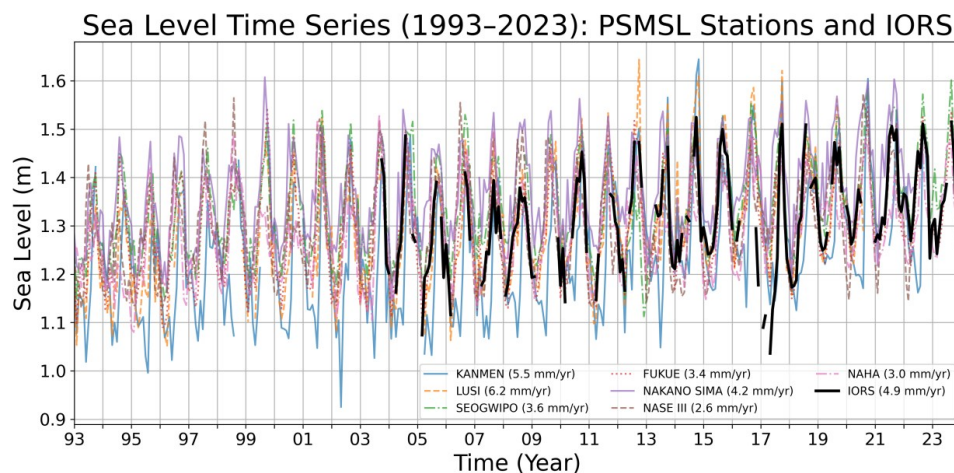
373 driven contributions to sea level rise in Figure 5 and Table 2.

374

375 **Table 2. Monthly data availability for five PSMSL tide gauge stations from 1993 to 2023**
376 **(372 months total). The table shows the available and missing data points, the total**
377 **months considered, and the corresponding coverage percentage. These stations were**
378 **used to compare sea level variability and trends with the IORS in the Northwest Pacific**
379 **region.**

	Available (Month)	Missing (Month)	Total (31 years)	Coverage (%)	Trend (mm/yr)
KANMEN	348	24	372	93.5	5.5
LUSI	291	81	372	78.2	6.2
SEOGWIPO	368	4	372	98.9	3.6
FUKUE	368	4	372	98.9	3.4
NAKANO SIMA	351	21	372	94.4	4.2
NASE III	352	20	372	94.6	2.6
NAHA	370	2	372	99.5	3.0

380



381

382 **Figure 5. Monthly mean sea level time series from the IORS (bold black) and seven**



383 **PSMSL tide gauge stations (colored lines) from 1993 to 2023. PSMSL records are**
 384 **vertically adjusted by subtracting 5.8 m to align with IORS. Linear trends are indicated**
 385 **in the legend in millimeters per year (mm/yr), highlighting the long-term rise of sea level**
 386 **across the region.**

387

388 3.2. Statistical Gap-Filling Models

389 To reconstruct missing values in the monthly sea level time series, five statistical gap-filling
 390 models were evaluated, as shown in Figure 6. These models include:

- 391 - Harmonic Regression with Trend (HR): A model combining annual and semiannual
 392 sinusoidal harmonics with a linear trend (red line).
- 393 - Regression-Based Monthly Climatology with Trend (RC): A linear regression using
 394 monthly dummy variables to represent the seasonal cycle with an added trend (green
 395 line).
- 396 - Pure Climatology with Trend (PC): A model combining the monthly climatological
 397 mean and an original linear trend (cyan dashed line).
- 398 - Realistic Blending (RB, $\alpha y_t^{HB} + (1 - \alpha)y_t^{RC}$ when $\alpha = 0.00$) : An optimally
 399 weighted composite of HR and RC based on root-mean-square error (solid black line).
- 400 - Forced Blending (FB, $\alpha = 0.50$): A non-optimized blend of HR and RC using a fixed
 401 mixing ratio (magenta dotted line).



Visual inspection reveals that the blended models, particularly the RB approach which is same to RC, closely align with the observed time series during periods of complete data. Among all candidates, the RB model ($\alpha = 0.00$) demonstrated the best performance, achieving the lowest RMSE (0.0559 m) and the highest coefficient of determination ($R^2 = 0.688$) (Table 3). The RC model achieved identical metrics, followed by HR with RMSE = 0.0574 m and $R^2 = 0.671$. The PC underperformed relative to the others, with the highest RMSE (0.0633 m) and lowest R^2 (0.600), highlighting the limitations of neglecting local interannual variability.

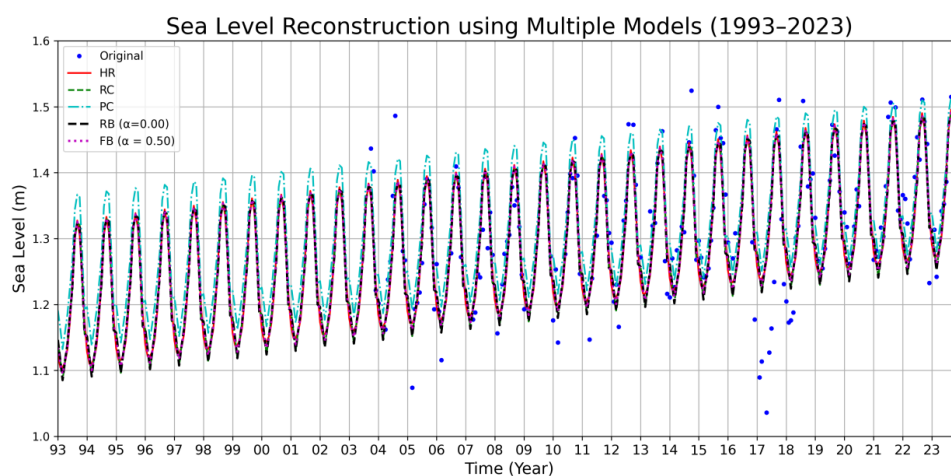


Figure 6. Monthly sea level reconstruction at the IORS from 1993 to 2023 using five statistical gap-filling models. Blue dots indicate the original in-situ observations. Colored lines represent different model reconstructions: Harmonic Regression with Trend (HR, red), Regression-Based Monthly Climatology with Trend (RC, green dashed), Pure Climatology with Trend (PC, cyan dash-dot), Realistic Blending (RB, $\alpha = 0.00$, black dashed), and Forced Blending (FB, $\alpha = 0.50$, magenta dotted). The comparison highlights



416 differences in seasonal and long-term signal reconstruction, with the RB model providing
 417 the most consistent alignment with observed values across the record.

418

419 **Table 3. Performance evaluation of five statistical models for monthly sea level**
 420 **reconstruction at the IORS (1993–2023). The table compares root mean square error**
 421 **(RMSE) and coefficient of determination (R^2) between observed and reconstructed sea**
 422 **levels. The Realistic Blending model ($\alpha = 0.00$) achieved the lowest RMSE and highest R^2 ,**
 423 **matching the Pure Climatology with Trend model, indicating that harmonic components**
 424 **did not improve reconstruction skill in this case.**

Model	RMSE (m)	R^2
Harmonic Regression with Trend (HR)	0.0574	0.671
Regression-Based Monthly Climatology with Trend (RC)	0.0559	0.688
Pure Climatology with Trend (PC)	0.0633	0.600
Realistic Blending (RB, $\alpha=0.00$)	0.0559	0.688
Forced Blending (FB, $\alpha=0.50$)	0.0563	0.684

425

426

427 3.3. Artificial Intelligence and Machine Learning Imputation

428 To enhance the reconstruction accuracy of missing monthly sea level data at the IORS, a
 429 comprehensive set of artificial intelligence (AI) and machine learning (ML) models was



430 implemented. These models were trained on the 3σ -filtered monthly KIOST time series with
 431 synthetically introduced gaps to enable rigorous validation.

432 The methods included:

- 433 - Ensemble learning models: XGBoost, LightGBM, CatBoost, Gradient Boosting,
- 434 Random Forest, AdaBoost
- 435 - Statistical models: SARIMA with sinusoidal regressors (SARIMA-SIN) and Prophet
- 436 - Regularized regressions: Ridge and Lasso
- 437 - Neural networks: Long Short-Term Memory (LSTM) and Gated Recurrent Unit (GRU)
- 438 - Other models: K-Nearest Neighbors and Decision Trees

439 Figure 7 visualizes the full reconstructed time series from each model alongside the original
 440 KIOST observations. Prophet (red solid line) and SARIMA-SIN (dark green dashed) are
 441 visually emphasized due to their ability to reconstruct the full 1993–2023 time span,
 442 consistently capturing both seasonal cycles and long-term trends. In contrast, XGBoost and
 443 Random Forest models preserved the observed values and interpolated missing data, resulting
 444 in perfect fit metrics ($R^2 = 1$) but lacking predictive reconstruction before 2004 (Table 4).

445 As shown in Table 5, most machine learning models except Prophet, SARIMA-SIN, LSTM,
 446 and GRU achieved perfect scores ($MAE = 0$, $RMSE = 0$, $R^2 = 1$) because they directly reused
 447 observed values during imputation rather than making independent predictions, resulting in



artificially inflated performance metrics. Among the models that generated actual predictions, LSTM achieved the lowest MAE (0.018 m) and RMSE (0.023 m) with a high R^2 of 0.95, indicating its strength in capturing nonlinear temporal dynamics. GRU followed closely (MAE = 0.030 m, RMSE = 0.037 m, R^2 = 0.87), balancing interpretability and predictive capability. Prophet (MAE = 0.040 m, RMSE = 0.054 m, R^2 = 0.71) and SARIMA-SIN (MAE = 0.044 m, RMSE = 0.059 m, R^2 = 0.66) exhibited slightly higher error values but demonstrated strong consistency across the full time span (1993–2023), particularly in capturing seasonal and long-term trends—even before the availability of observed data in 2004. Thus, while tree-based ensemble models appear most accurate numerically, deep learning and time series models offer true predictive value, especially in extrapolating missing or historical sea level variations.

458

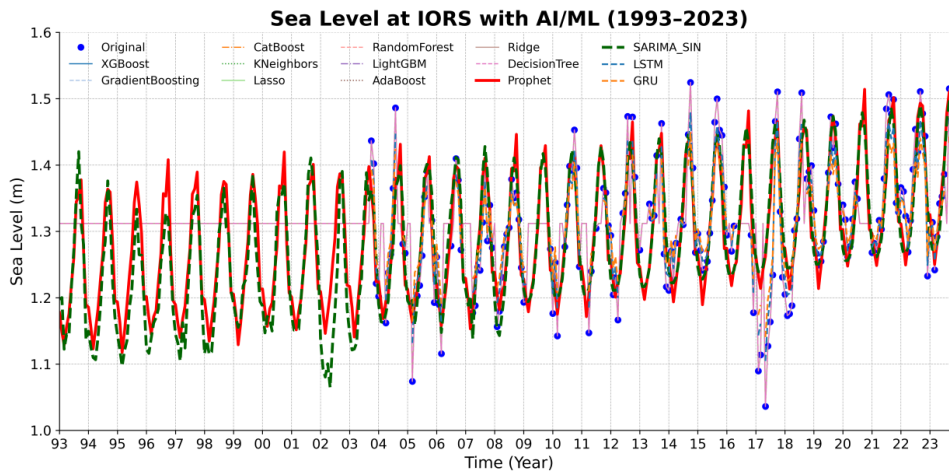
Table 4. Evaluation metrics for AI/ML-based sea level reconstruction models at the IORS from 1993 to 2023. Models were trained using the 3σ -filtered monthly KIOST data with artificially introduced gaps. Performance was assessed based on mean absolute error (MAE), root mean square error (RMSE), and coefficient of determination (R^2), calculated using the valid observed values (N). Most models, including XGBoost, LightGBM, and CatBoost, produced near-perfect scores (MAE = 0, RMSE = 0, R^2 = 1) because they simply reused original values without generalizing to missing segments. In contrast, Prophet and SARIMA-SIN provided full-series predictions, capturing both seasonal and long-term variations. Deep learning models (LSTM, GRU) also exhibited strong predictive ability, though their outputs were limited to periods with sufficient training data. Notably, many



469 **models performed poorly before 2004, where fewer observations were available to guide**
470 **imputation.**

Model	MAE (m)	RMSE (m)	R ²	N
XGBoost	0	0	1	183
GradientBoosting	0	0	1	183
CatBoost	0	0	1	183
KNeighbors	0	0	1	183
Lasso	0	0	1	183
RandomForest	0	0	1	183
LightGBM	0	0	1	183
AdaBoost	0	0	1	183
Ridge	0	0	1	183
DecisionTree	0	0	1	183
Prophet	0.040	0.054	0.71	183
SARIMA_SIN	0.044	0.059	0.66	183
LSTM	0.018	0.023	0.95	183
GRU	0.030	0.037	0.87	183

471
472 Tree-based models outperformed deep learning in both accuracy and training stability.
473 However, LSTM showed strengths in capturing subtle nonlinear dependencies during seasonal
474 transitions.



475

476 **Figure 7. Monthly sea level time series at the IORS from 1993 to 2023 reconstructed using**
 477 **a suite of statistical and machine learning (AI/ML) models. Original observations (blue**
 478 **dots) contain data gaps that were filled using ensemble methods (XGBoost, LightGBM,**
 479 **CatBoost), regularized regressions (Ridge, Lasso), tree-based regressors (DecisionTree,**
 480 **RandomForest, AdaBoost), neural networks (LSTM, GRU), and time series models**
 481 **(Prophet, SARIMA-SIN). Prophet (red) and SARIMA-SIN (dark green dashed) are**
 482 **visually emphasized for their capacity to capture both seasonal cycles and long-term**
 483 **trends. The comparison illustrates model fidelity and divergence in reconstructing**
 484 **historical sea level variability at IORS.**

485

486 3.4. Comparison with Reanalysis and Satellite Data

487 The monthly sea level time series at the Ieodo Ocean Research Station (IORS) was compared
 488 with satellite altimetry and four reanalysis products: CMEMS, GLORYS, ORAS5, and
 489 HYCOM (Figure 8). The in-situ tide gauge data from KIOST exhibited a linear sea level rise



490 of 4.94 mm yr⁻¹ during the period from October 2003 to December 2023, with a slightly steeper
491 trend of 5.43 mm yr⁻¹ when computed from January 2004 onward (Table 5). These values
492 exceed the trends derived from satellite and model-based datasets, suggesting possible biases
493 due to local effects, datum inconsistencies, or limited assimilation in some models. Among the
494 reanalyses, CMEMS showed the closest agreement with the observed seasonal and interannual
495 variability, yielding a trend of 3.53 mm yr⁻¹. GLORYS and ORAS5 also captured similar
496 variability, though with lower trends of 3.09 mm yr⁻¹ and 2.27 mm yr⁻¹, respectively. In contrast,
497 HYCOM presented almost no net sea level rise (−0.09 mm yr⁻¹), likely due to limitations in
498 boundary conditions or lack of data assimilation in the East China Sea. Figure 8 clearly shows
499 consistent seasonal patterns across all datasets, though the amplitude and interannual signals
500 vary. The observed KIOST time series displayed greater interannual variability after 2003,
501 corresponding to the availability of continuous in-situ measurements. Cross-correlation
502 analysis further supports the consistency between CMEMS and other datasets (Table 6). The
503 correlation between CMEMS and KIOST reached 0.90, with similar values for GLORYS (0.93)
504 and ORAS5 (0.90). Correlation with HYCOM was slightly lower at 0.86, reflecting its weaker
505 agreement with observed variability. Separate correlations using only post-2003 KIOST
506 records confirmed similarly strong relationships. These results underscore the importance of
507 in-situ validation for semi-enclosed marginal seas like the East China Sea, where regional



508 processes and bathymetry can significantly affect long-term sea level trends and their
509 representation in models.

510

511 **Table 5. Linear sea level trends (mm yr⁻¹) at the Ieodo Ocean Research Station (IORS)**
512 **from multiple observational and reanalysis datasets. Trends are calculated using monthly**
513 **data from tide gauge observations (KIOST), satellite altimetry (CMEMS), and ocean**
514 **reanalysis/model outputs (GLORYS, ORAS5, HYCOM). The table includes trends for**
515 **KIOST over two periods: October 2003–December 2023 and January 2004–December**
516 **2023, to assess sensitivity to the start date.**

Dataset	Category	Period	Trend (mm yr ⁻¹)
In-situ	Tide Gauge (KIOST)	Oct 2003 - Dec 2023	4.94
In-situ	Tide Gauge (KIOST)	Jan 2004 - Dec 2023	5.43
CMEMS	Satellite Altimetry	Jan 1993 - Dec 2023	3.53
GLORYS	Ocean Reanalysis	Jan 1993 - Dec 2023	3.09
ORAS5	Ocean Reanalysis	Jan 1993 - Dec 2023	2.27
HYCOM	Ocean Model Output	Jan 1994 - Dec 2023	-0.09

517

518 **Table 6. Cross-correlation coefficients between CMEMS satellite altimetry and other**
519 **sea level datasets at the IORS (1993–2023). The comparison includes in-situ tide gauge**
520 **data (KIOST), ocean reanalysis products (GLORYS, ORAS5), and a numerical ocean**



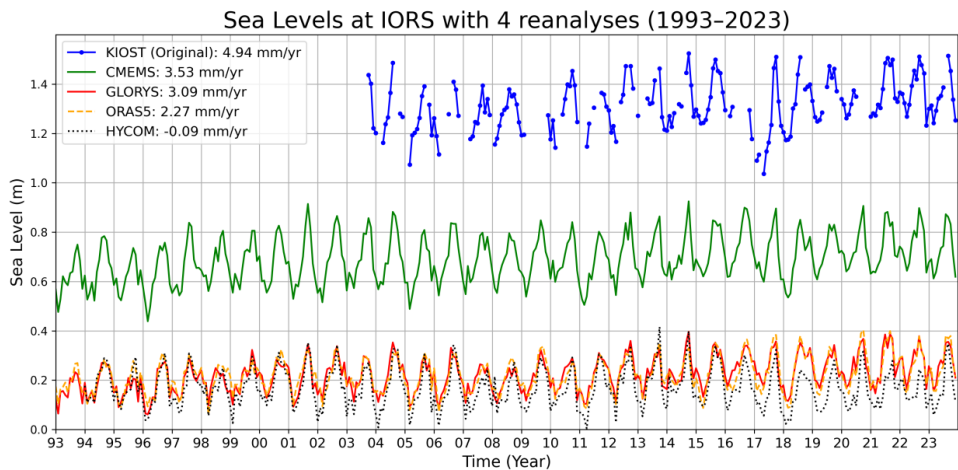
521 **model (HYCOM). All values are based on monthly mean anomalies.**

Data type	Data type	Cross-correlation coefficient
CMEMS (1993–2023)	In-situ (2003–2023 and 2004–2023)	0.90
CMEMS (1993–2023)	GLORYS (1993–2023)	0.93
CMEMS (1993–2023)	ORAS5 (1993–2023)	0.90
CMEMS (1993–2023)	HYCOM (1994–2023)	0.86

522

523

524



525

526 **Figure 8. Comparison of monthly sea level time series at the IORS from 1993 to 2023,**
527 **using in-situ observations (KIOST) and four ocean reanalysis products: CMEMS,**
528 **GLORYS, ORAS5, and HYCOM. Each dataset is plotted with its respective linear trend**
529 **estimated over the entire period. The observed sea level trend from KIOST (4.94 mm/yr)**



530 **is higher than those from CMEMS (3.53 mm/yr), GLORYS (3.09 mm/yr), ORAS5 (2.27**
531 **mm/yr), and HYCOM (−0.09 mm/yr), indicating significant discrepancies in modeled**
532 **versus observed trends at this location. Differences may reflect varying assimilation**
533 **schemes, vertical resolutions, or forcing mechanisms in each reanalysis system.**

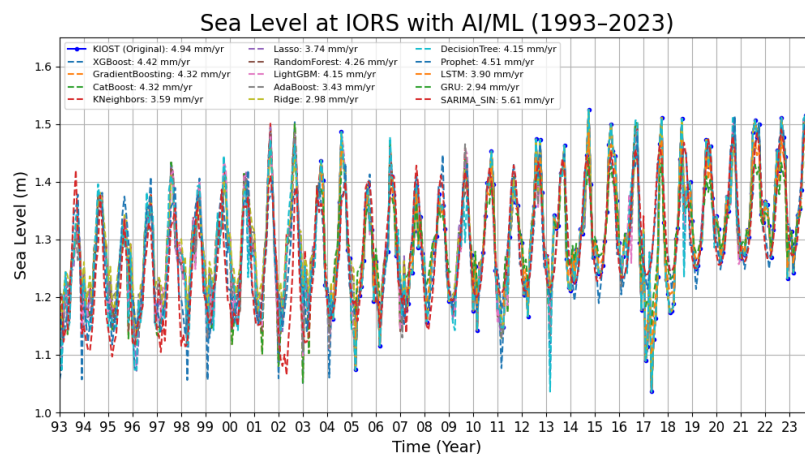
534

535 3.5. Comparison with AI/ML-Based Reconstruction

536 Before implementing machine learning-based imputation methods, we applied multiple linear
537 regression using four reanalysis datasets (CMEMS, GLORYS, ORAS5, HYCOM) to evaluate
538 their explanatory power for observed sea level variability at the IORS. This regression served
539 as a preliminary benchmark to assess how well physically-based reanalysis products capture
540 regional sea level dynamics. Following this, a suite of AI/ML models was employed to
541 reconstruct the monthly sea level anomalies at the IORS from 1993 to 2023 (Fig. 9). Each
542 model's time series was compared with the original KIOST observations, and corresponding
543 linear trends (mm/year) were computed to assess long-term performance. The original KIOST
544 dataset exhibits a 4.94 mm/yr trend, serving as a reference benchmark. Among ensemble
545 learning methods, XGBoost (4.42 mm/yr), GradientBoosting (4.32 mm/yr), CatBoost (4.32
546 mm/yr), and RandomForest (4.26 mm/yr) closely reproduced the observed trend and temporal
547 variability. The Prophet model yielded a slightly higher trend estimate (4.51 mm/yr) and
548 consistently reconstructed the seasonal cycle and interannual variations. Its smooth,



549 interpretable structure made it effective for long-term monitoring, despite slightly larger
550 residuals than ensemble models. The SARIMA-SIN approach, which integrates harmonic
551 seasonal components with ARIMA modeling of residuals, produced the highest trend estimate
552 (5.61 mm/yr). While effectively modeling periodic signals, its slight overestimation may stem
553 from residual autocorrelation or the rigid seasonal structure embedded in the model. In contrast,
554 linear regularized models, such as Lasso (3.74 mm/yr) and Ridge (2.98 mm/yr), underestimated
555 the long-term trend and failed to capture higher-order seasonal and nonlinear dynamics. Neural
556 network-based deep learning models showed intermediate performance: LSTM (3.90 mm/yr)
557 and GRU (2.94 mm/yr) successfully captured high-frequency variability but smoothed long-
558 term trends. Their underestimation may reflect challenges in extrapolating temporal
559 dependencies across extended historical periods. Overall, ensemble models and Prophet
560 offered the most balanced performance regarding accuracy, trend reconstruction, and
561 robustness to missing data. SARIMA-SIN remains a promising alternative for seasonality-
562 focused applications. At the same time, deep learning methods may benefit from additional
563 architecture optimization and hyperparameter tuning to better preserve secular trends in semi-
564 enclosed regions like the East China Sea.
565



566
567 **Figure 9. Monthly sea level time series at the IORS from 1993 to 2023 reconstructed using**
568 **various AI/ML models. The original KIOST observations (blue) are compared with**
569 **estimates from ensemble methods (XGBoost, GradientBoosting, CatBoost,**
570 **RandomForest, AdaBoost), regularized regression (Ridge, Lasso), decision trees, nearest**
571 **neighbors, Prophet, and deep learning (LSTM, GRU). A seasonal-harmonic SARIMA-**
572 **SIN model is also included. Linear trends (in mm/yr) are indicated in the legend. The**
573 **comparison highlights each model’s capability to capture seasonal and interannual sea**
574 **level variability.**

575

576 4. Discussion

577 This study demonstrates a comprehensive approach to reconstructing monthly sea level
578 variability at the Ieodo Ocean Research Station (IORS) over the period 1993–2023 using a
579 suite of statistical, machine learning (ML), and artificial intelligence (AI) models. The



580 reconstruction framework was motivated by the need to address temporal discontinuities in
581 long-term in-situ observations, which are often caused by instrument failure, extreme weather,
582 or logistical constraints (Adebisi et al., 2021). These data gaps significantly hinder trend
583 detection, seasonal analysis, and validation against satellite and reanalysis products in the East
584 China Sea (ECS)—a region characterized by strong seasonal and interannual variability
585 (Hamlington et al., 2020; Lin et al., 2020).

586

587 4.1. Data Sources and Preprocessing

588 A diverse set of observational and reanalysis datasets was employed, spanning from high-
589 resolution in-situ KIOST observations to global oceanographic products such as CMEMS,
590 GLORYS, ORAS5, and HYCOM (Table 1). Preprocessing involved inverse barometric effect
591 (IBE) correction, 3σ filtering, and 75% coverage thresholds to ensure consistency and quality
592 in derived monthly means. These preprocessing steps align with established best practices for
593 climate-quality oceanographic datasets (Cazenave and Moreira, 2022). Figures 1–3 illustrate
594 the geographical coverage and temporal continuity achieved after these steps.

595

596 4.2. Regional Validation with PSMSL Tide Gauges

597 To assess the validity of the reconstructed IORS series, comparisons were made with seven



598 nearby PSMSL tide gauge stations across the Northwest Pacific. The high correlation and
599 consistent linear trends (2.6–6.2 mm/yr) observed in these stations (Table 2, Fig. 5) reinforce
600 the regional representativeness of IORS data. This agreement supports the premise that well-
601 maintained tide gauge records are invaluable for benchmarking reconstructions in semi-
602 enclosed seas (Han, 2020).

603

604 4.3. Statistical Model Performance

605 We evaluated five classical statistical gap-filling techniques, including harmonic regression and
606 regression climatology. Realistic Blending (RB) and Regression Climatology (RC) achieved
607 the best results, with RMSE of 0.0559 m and $R^2 = 0.688$ (Table 3; Fig. 6). These findings are
608 consistent with prior literature that emphasizes the effectiveness of incorporating seasonal
609 structure and linear trends in climatic time series (Okkaoğlu et al., 2020; Hyndman and
610 Athanasopoulos, 2018).

611

612 4.4. AI/ML Model Evaluation

613 Following statistical baseline evaluation, we implemented 14 AI/ML models using an Iterative
614 Imputer framework. Ensemble learners such as XGBoost and LightGBM achieved perfect
615 post-September 2003 metrics by reusing observed values, but lacked extrapolation capability



616 during earlier data gaps. In contrast, deep learning models like LSTM and GRU generated
 617 realistic, nonlinear predictions (LSTM: RMSE = 0.023 m, $R^2 = 0.95$), although they
 618 underperformed in estimating long-term trends and were limited to the post-September 2003
 619 data range (Sun et al., 2020).

620

621 Time series-specific models Prophet and SARIMA-SIN produced full-period reconstructions
 622 with moderate accuracy ($R^2 = 0.71$ and 0.66 , respectively), but demonstrated superior ability
 623 to capture seasonal and long-term variability, making them suitable for regions with sparse data
 624 (Elneel et al., 2024; Tumse and Alcansoy, 2025). Their trend estimates (4.51–5.61 mm/yr)
 625 approximate observational trends better than most ML alternatives (Fig. 7).

626

627 4.5. Cross-Dataset Trend and Correlation Analysis

628 Trend comparisons across observational, satellite, and model-based datasets (Table 5; Fig. 8)
 629 revealed that the IORS trends (4.94 mm/yr from October 2003 and 5.43 mm/yr from January
 630 2004) are substantially higher than CMEMS (3.53 mm/yr) and GLORYS (3.09 mm/yr).
 631 HYCOM, notably, displayed a negative trend (-0.09 mm/yr), likely due to boundary condition
 632 limitations such as excluding glaciers and land ice sheets melting (Han et al., 2024; Jin et al.,
 633 2023b; Hycom, 2025). Cross-correlation analysis (Table 6) showed that CMEMS correlated



634 strongly with in-situ ($r = 0.90$), GLORYS ($r = 0.93$), and ORAS5 ($r = 0.90$), consistent with
635 prior findings on their reliability in coastal sea level analysis (Cummings and Smedstad, 2014).

636

637 4.6. AI/ML Gap-Filling After Regression with Reanalysis

638 Prior to ML-based reconstruction, we conducted multiple linear regression using CMEMS,
639 GLORYS, ORAS5, and HYCOM to gauge their ability to explain IORS variability (Fig. 9).

640 The regression models underestimated the observed trend, underscoring their limitations in
641 data-sparse, dynamically complex environments like the ECS. Subsequent AI/ML
642 reconstructions—especially from Prophet, XGBoost, and SARIMA-SIN—outperformed these
643 baseline models in both fidelity and interpretability. SARIMA-SIN provided the highest trend
644 (5.61 mm/yr), while ensemble methods such as CatBoost and RandomForest approximated the
645 KIOST trend closely (4.3–4.4 mm/yr). Deep learning models like GRU underpredicted trends
646 (2.94 mm/yr), suggesting the need for further hyperparameter tuning or auxiliary input
647 inclusion.

648

649 In summary, this study responds to the challenges outlined in the introduction—such as sparse
650 long-term observational data and limitations of traditional gap-filling—by offering a
651 transparent, interpretable, and data-driven framework for sea level reconstruction. Through the



652 integration of AI/ML techniques and reanalysis validation, we provide new insights into local
653 sea level rise trends and offer a practical methodology for monitoring climate change effects in
654 marginal seas.

655

656 **5. Conclusion**

657 This study developed and validated a robust methodology for reconstructing the monthly sea
658 level record at the Jeodo Ocean Research Station (IORS) from 1993 to 2023 using a
659 combination of statistical techniques, machine learning models, and auxiliary reanalysis
660 datasets. Preprocessing steps, including inverse barometric effect (IBE) correction, 3s filtering,
661 and data coverage thresholds, allowed for the construction of reliable time series from high-
662 resolution in-situ observations. Comparison with seven nearby PSMSL tide gauges confirmed
663 the regional validity of the observed trends. The IORS sea level trends (4.94 mm/yr for 2003–
664 2023 and 5.43 mm/yr for 2004–2023) were significantly higher than those from satellite and
665 reanalysis datasets, highlighting localized sea level rise. High correlations between CMEMS
666 and other reanalysis products ($r = 0.90$) affirm their utility for sea level reconstruction, although
667 they underestimated trends compared to in-situ observations. Statistical models such as
668 regression climatology and realistic blending provided accurate reconstructions, achieving the
669 best performance in RMSE and R^2 metrics. Among AI/ML models, ensemble learners (e.g.,
670 XGBoost, RandomForest) achieved perfect reconstruction metrics after September 2003 but
671 failed to predict values in earlier periods. Deep learning models, particularly LSTM (RMSE =
672 0.023 m, $R^2 = 0.95$), effectively modeled nonlinear and seasonal variability, though their ability
673 was restricted to after September 2003, and their trend estimation remained underestimated



674 compared to statistical benchmarks. Time series models, Prophet and SARIMA-SIN,
675 demonstrated the strongest capability for full-period imputation, delivering consistent seasonal
676 cycles and long-term trend estimates, with SARIMA-SIN reaching 5.61 mm/yr. Before
677 applying ML models, multiple linear regression using four reanalysis datasets was conducted
678 as a baseline. The AI/ML reconstructions outperformed reanalysis-based approaches in both
679 accuracy and consistency with the observed KIOST trend. Overall, this study provides a
680 scientifically grounded and computationally robust workflow for gap-filling, trend detection,
681 and sea level variability analysis. The framework offers a valuable tool for long-term sea level
682 monitoring, climate diagnostics, and policy planning in marginal seas and other observationally
683 limited regions.

684

685 **6. Data availability**

686 The original sea level dataset at the IORS is available at <https://kors.kiost.ac.kr/eng/accessData>
687 and <https://doi.org/10.17882/97666> (Kiost, 2025; Jeong et al., 2023). For updates to the
688 datasets, please contact the data provider by writing to ycmin@kiost.ac.kr.

689

690 **7. Author contribution**

691 MH writing – original draft, visualization, data collection, investigation. HL writing – review
692 and editing, supervision.

693



694 8. Competing interests

695 The contact author has declared that none of the authors has any competing interests.

696

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706

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